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# Numerical cell model investigating cellular carbon fluxes in *Emiliania huxleyi*



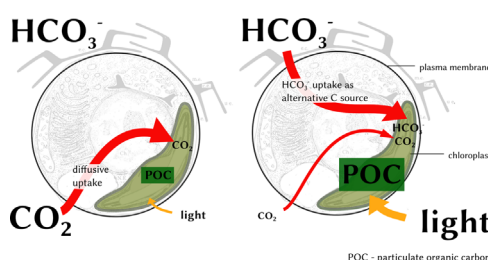
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## HIGHLIGHTS

- Cellular C fluxes through coccolithophores are poorly understood.
- By means of a numerical cell model, we investigate cellular C fluxes through *Emiliania huxleyi*.
- In the model,  $\text{CO}_2$  and  $\text{HCO}_3^-$  rarely inter-convert within the cytosol.
- At high external  $[\text{CO}_2]$  and low light, photosynthetic C demand can be covered by diffusive  $\text{CO}_2$  uptake.
- $\text{CO}_2$  accumulation around RubisCO is possible without transporting C species uphill.

## GRAPHICAL ABSTRACT



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## ABSTRACT

Coccolithophores play a crucial role in the marine carbon cycle and thus it is interesting to know how they will respond to climate change. After several decades of research the interplay between intracellular processes and the marine carbonate system is still not well understood. On the basis of experimental findings given in literature, a numerical cell model is developed that describes inorganic carbon fluxes between seawater and the intracellular sites of calcite precipitation and photosynthetic carbon fixation. The implemented cell model consists of four compartments, for each of which the carbonate system is resolved individually. The four compartments are connected to each other via  $\text{H}^+$ ,  $\text{CO}_2$ , and  $\text{HCO}_3^-$  fluxes across the compartment-confining membranes. For  $\text{CO}_2$  accumulation around RubisCO, an energy-efficient carbon concentrating mechanism is proposed that relies on diffusive  $\text{CO}_2$  uptake. At low external  $\text{CO}_2$  concentrations and high light intensities,  $\text{CO}_2$  diffusion does not suffice to cover the carbon demand of photosynthesis and an additional uptake of external  $\text{HCO}_3^-$  becomes essential. The model is constrained by data of *Emiliania huxleyi*, the numerically most abundant coccolithophore species in the present-day ocean.

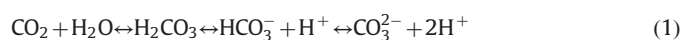
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## 1. Introduction

Human activity has rapidly increased atmospheric carbon dioxide partial pressure ( $\text{pCO}_2$ ) from about 280  $\mu\text{atm}$  in year 1750 to around 400  $\mu\text{atm}$  in 2013 (NOAA, 2013). A continuing strong increase is expected. Carbon dioxide ( $\text{CO}_2$ ) dissolves in seawater and becomes part of the ocean's carbonate system, whose simplified

representation is:



Due to the inter-conversion between individual inorganic carbon species (iCS) and the high total alkalinity of seawater, the ocean exhibits a strong capacity to buffer anthropogenic  $\text{CO}_2$  emissions. The uptake of  $\text{CO}_2$ , however, impacts the ocean's carbonate chemistry with potentially tremendous consequences for marine organisms and whole ecosystems. In this study, we focus on coccolithophores, unicellular calcareous algae that belong to the main pelagic calcifiers.

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## Nomenclature

### Abbreviations used in text

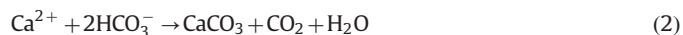
$\Omega$	Calcite saturation product
CA	Carbonic anhydrase
CCM	CO <sub>2</sub> /carbon concentrating mechanism
CV	Calcification/coccolith vesicle
DIC	Dissolved inorganic carbon, i.e. sum of CO <sub>2</sub> , H <sub>2</sub> CO <sub>3</sub> , HCO <sub>3</sub> <sup>-</sup> , and CO <sub>3</sub> <sup>2-</sup>
HC	High carbon

HL	High light
iCS	Inorganic carbon species, here CO <sub>2</sub> , HCO <sub>3</sub> <sup>-</sup> , or CO <sub>3</sub> <sup>2-</sup>
LC	Low carbon
LL	Low light
MIMS	Membrane inlet mass spectrometry
pCO <sub>2</sub>	CO <sub>2</sub> partial pressure
PIC	Particulate inorganic carbon, here calcite
POC	Particulate organic carbon
RubisCO	Ribulose-1,5-bisphosphate-Carboxylase-Oxygenase

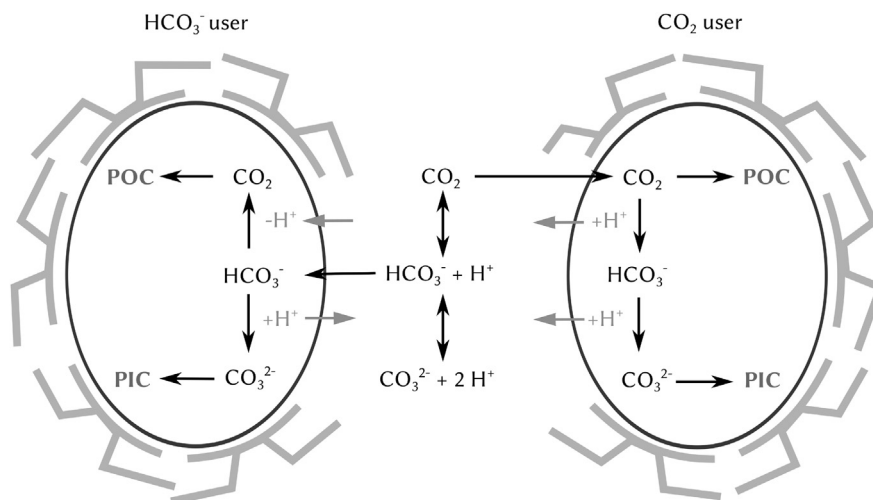
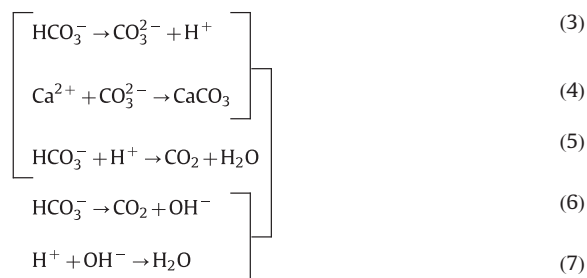
Most notably, coccolithophores impact the ocean's carbon cycle via the formation of particulate organic and inorganic carbon (POC and PIC). Our model organism *Emiliania huxleyi* is the most abundant and one of the best studied coccolithophore species.

During photosynthesis, POC is built from CO<sub>2</sub> inside the chloroplast. Particulate inorganic carbon is precipitated in the form of calcium carbonate (CaCO<sub>3</sub>) from carbonate ions (CO<sub>3</sub><sup>2-</sup>) inside a specialised cell compartment termed calcification or coccolith vesicle (CV). Since internal pH homeostasis and charge neutrality have to be ensured (Wolf-Gladrow et al., 2007), the production of one mole POC either relies on one mole external CO<sub>2</sub> or one mole of bicarbonate ions (HCO<sub>3</sub><sup>-</sup>) plus one mole of protons (H<sup>+</sup>) (cf. Fig. 1). Correspondingly, the production of one mole PIC (i.e. CaCO<sub>3</sub>) relies on one mole Ca<sup>2+</sup> plus either one mole external HCO<sub>3</sub><sup>-</sup> and the extrusion of one mole H<sup>+</sup> or one mole CO<sub>2</sub> and the extrusion of two moles H<sup>+</sup>. The productions of POC and PIC thus impact the marine carbonate system differently: Particulate organic carbon production, i.e. CO<sub>2</sub> consumption, reduces the DIC (dissolved inorganic carbon) concentration in seawater and shifts the carbonate system towards higher pH values. Particulate inorganic carbon production, i.e. CO<sub>3</sub><sup>2-</sup> consumption, decreases DIC and alkalinity in the molar ratio 1:2 and shifts the carbonate system towards lower pH values. In view of bulk seawater, it does not make a difference whether carbon for POC and PIC production is taken up in the form of CO<sub>2</sub> or HCO<sub>3</sub><sup>-</sup>. The composition of the external (seawater) carbonate system, however, impacts uptake rates of external iCS into the cell and therewith the fixation rates of internal iCS into POC and PIC. Intracellular pathways and conversion rates of iCS cannot be measured directly and are thus not known. Therewith, it is difficult to understand the interdependency between the marine carbonate system and POC and PIC production rates mechanistically.

In order to reveal the influence of abiotic conditions on the PIC: POC ratio of coccolithophores in a systematic manner, intracellular pathways of iCS have to be elucidated first. As mentioned above, it is not possible to measure intracellular iCS fluxes such as the CO<sub>2</sub> flux across the chloroplast envelope directly and *in vivo*. We hence rely on indirect information about these fluxes. External carbon sources of POC and PIC production have been an issue of discussion since more than five decades. Paasche (1964) demonstrated by means of short-term <sup>14</sup>C incubations that external HCO<sub>3</sub><sup>-</sup> is the main substrate for calcite production in *E. huxleyi*. The bulk of experimental work that was conducted afterwards has led to the same conclusion (Berry et al., 2002; Rost et al., 2002). For *E. huxleyi*, Paasche (1964) suggested that CO<sub>2</sub> intracellularly produced in the reaction



might be used in photosynthesis. Reaction (2) thus presumes external HCO<sub>3</sub><sup>-</sup> being the carbon source for both, photosynthesis and calcite precipitation, and gives the sum of reactions (3) to (5) or (3), (4), (6), and (7) (cf. Zeebe and Wolf-Gladrow, 2001).



**Fig. 1.** Theoretical H<sup>+</sup> budget related to the intracellular conversion of inorganic carbon species for a HCO<sub>3</sub><sup>-</sup> using cell (left) and a CO<sub>2</sub> using cell (right). Photosynthesis and calcite precipitation are the two major processes that require dissolved inorganic carbon. PIC - particulate inorganic carbon, POC - particulate organic carbon.

In comparison to the other three reactions, reactions (5) and (6) are slow. Reaction (5) can be accelerated by the enzyme carbonic anhydrase (CA) and is favoured under low pH values. Reactions (3), (4), (6), and (7) occur under high pH values and may hence constitute prominent reactions inside the CV. The stoichiometry, however, that is implied by reaction (2), i.e. release of one mole  $\text{CO}_2$  per mole precipitated  $\text{CaCO}_3$ , is unlikely to occur, primarily because the reaction rate of reaction (3) exceeds the one of reaction (6) by more than five orders of magnitude (at temperature  $15^\circ\text{C}$  and salinity 30). Nevertheless, the  $\text{H}^+$  released during calcite precipitation (eqns. (3) and (4)) could contribute to  $\text{CO}_2$  formation for photosynthetic use in another compartment (eqn. (5)). The relative strength of this coupling process would on the one hand depend on the sum of intracellular  $\text{H}^+$  fluxes and on the other hand on the external carbon source for photosynthesis. In the same work, Paasche (1964) showed that external  $\text{HCO}_3^-$  as well as external  $\text{CO}_2$  are used for photosynthesis. Later on, many experiments were conducted aiming to determine the external carbon source for photosynthesis. They partly lead to contradicting results (for an overview until year 2001 see Paasche, 2002). Rost et al. (2002) concluded from the distribution of stable carbon isotopes in POC that  $\text{CO}_2$  as well as  $\text{HCO}_3^-$  are used for photosynthesis. Based on measurements of  $^{14}\text{C}$  fluxes in disequilibrium, Rokitta and Rost (2012) identified external  $\text{HCO}_3^-$  as the major iCS source that supports POC production. The  $^{14}\text{C}$  disequilibrium assay they used, however, was conducted at a higher pH value than the cells were acclimated to. Kottmeier et al. refined this  $^{14}\text{C}$  disequilibrium assay, which enabled them to measure the external iCS source of POC at different assay pH values. They found external  $\text{CO}_2$  being the major source for photosynthesis under current  $\text{CO}_2$  conditions.

To sum up, it seems that external  $\text{CO}_2$  and  $\text{HCO}_3^-$  are the carbon sources for POC production and external  $\text{HCO}_3^-$  is the only carbon source for PIC production which suggests that conversion of  $\text{CO}_2$  to  $\text{HCO}_3^-$  inside the cytosol is low.

On basis of current knowledge about membrane transporters and enzymes in coccolithophores as well as in phytoplankton and plant cells in general, a numerical cell model is developed to investigate the pathways of inorganic carbon from the external medium towards the sites of calcite precipitation and photosynthesis. The model is used to examine the following three questions:

1. Can  $\text{CO}_2$  and  $\text{HCO}_3^-$  pass the cytosol without being inter-converted?
2. Can enough external  $\text{CO}_2$  be taken up by diffusion to satisfy the carbon requirement of POC production?

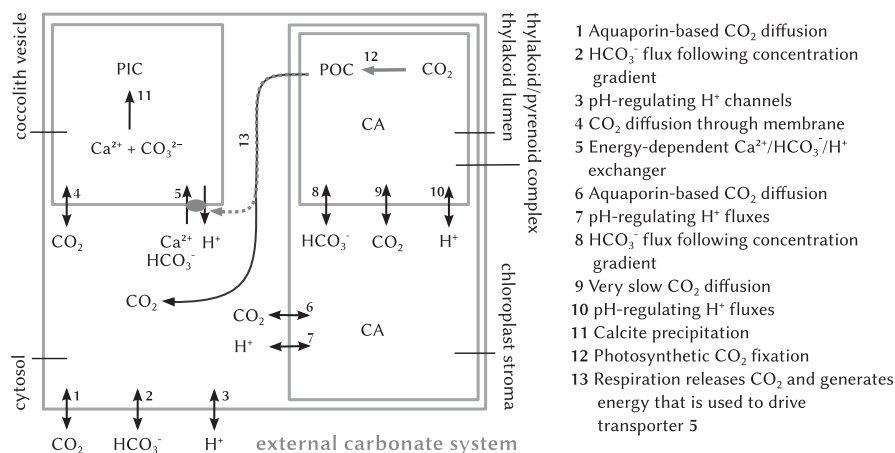
3. Can  $\text{CO}_2$  be accumulated around RubisCO without inorganic carbon species being transported against concentration gradients?

## 2. Cell model

The cell is divided into four compartments: The cytosol, the CV, the chloroplast stroma, and the thylakoid/pyrenoid complex (Fig. 2). Our working hypothesis concerning the carbon concentrating mechanism (CCM), i.e. the mechanism that leads to  $\text{CO}_2$  accumulation around RubisCO, is illustrated in Figure 3. Membrane- and compartment-specific characteristics assumed for the model's set-up are briefly described in the following. In Section 2.1, the mathematical implementation is described.

**Plasma membrane.** Aquaporins inside the plasma membrane facilitate diffusive  $\text{CO}_2$  uptake (1 in Fig. 2). We use the permeability coefficient for  $\text{CO}_2$  diffusion through the plasma membrane of *Nicotiana tabacum* which contains aquaporins (Uehlein et al., 2008). Protons pass the plasma membrane through voltage-gated  $\text{H}^+$  channels in coccolithophores (passive direction: out of the cell) (Taylor et al., 2011). In the model, cytosolic pH (ca. 7; Anning et al., 1996) is regulated by means of these channels (3 in Fig. 2). Anion exchangers that are associated with  $\text{HCO}_3^-$  fluxes (assumingly across the plasma membrane) were found in *E. huxleyi* (Herfort et al., 2002; von Dassow et al., 2009; Mackinder et al., 2011; Taylor et al., 2011; Rokitta and Rost, 2012; Bach et al., 2013). In our model, external  $\text{HCO}_3^-$  enters the cytosol following its concentration gradient (2 in Fig. 2) and thus cannot accumulate inside the cytosol. *In vivo*,  $\text{HCO}_3^-$  probably enters the cytosol in sym- or antiport with or against other ions. One possibility is a cotransport with  $\text{Na}^+$  that may be energised by the electrochemical gradient driving  $\text{Na}^+$  into the cell: The inner face of the plasma membrane is charged negatively and the external concentration of  $\text{Na}^+$  furthermore exceeds the one inside the cytosol. As the internal  $\text{Na}^+$  level has to remain low,  $\text{Na}^+$  would subsequently have to be exported (potentially in antiport against  $\text{H}^+$ ) at the expense of energy equivalents.

**Coccolith vesicle membrane.** We assume that  $\text{CO}_2$  diffuses across the CV membrane directly, i.e. not through aquaporins (4 in Fig. 2), and use the  $\text{CO}_2$  permeability coefficient measured for membranes of pure liposomes (Prasad et al., 1998). In the model, energy requiring  $\text{Ca}^{2+}/\text{HCO}_3^-/\text{H}^+$  transporters (5 in Fig. 2, cf. Holtz et al., 2013) import the substrates for calcite precipitation against their concentration gradients and support the establishment of an



**Fig. 2.** Cell model including four compartments. Fluxes are indicated by arrows. Grey coloured arrows indicate rates that are set to a predefined value. The external carbonate system which forces cellular in- and efflux rates is set to the values measured by Rokitta and Rost (2012). CA - carbonic anhydrase; PIC - particulate inorganic carbon; POC - particulate organic carbon.

alkaline milieu that is favourable for calcite precipitation. These transporters are not known from literature. The assumed stoichiometry of  $1 \text{ Ca}^{2+} : 1 \text{ HCO}_3^- : 1 \text{ H}^+$ , however, ensures charge balance and also sufficiently high calcite saturation states ( $\Omega$ ) (cf. Wolf-Gladrow et al., 2007; Holtz et al., 2013). Calcite precipitation can take place at  $\Omega > 1$  and it increases with increasing  $\Omega$  values (11 in Fig. 2).

**Chloroplast envelope and stroma.** The chloroplast of coccolithophores is surrounded by four membranes, one of which does not constitute a significant barrier for  $\text{CO}_2$  diffusion because it includes large pores. The implemented chloroplast stroma is thus surrounded by three membranes only which are taken into consideration via a  $\text{CO}_2$  permeability coefficient that is three times lower than the one of a single membrane layer (cf. Thoms et al., 2001). We assume that all three membranes contain aquaporins to facilitate diffusive influx of  $\text{CO}_2$  into the chloroplast (6 in Fig. 2). The existence of aquaporins in the chloroplast envelope has been shown for *N. tabacum* (Uehlein et al., 2008). The pH of the *in silico* chloroplast stroma (ca. 8) can be regulated via proton fluxes across the chloroplast envelope (7 in Fig. 2). Following the findings of Quiroga and González (1993) for coccolithophore CCMP-299, we assume CA activity inside the chloroplast stroma (outside the pyrenoid).

**Thylakoid membrane and lumen.** We assume that  $\text{HCO}_3^-$ ,  $\text{H}^+$ , and  $\text{CO}_2$  fluxes connect the thylakoid lumen to the chloroplast stroma (8, 9, and 10 in Fig. 2). *In vivo*,  $\text{HCO}_3^-$  could enter the thylakoid through channels driven by the  $\text{HCO}_3^-$  concentration gradient and the positive charging of the thylakoids' inner face (due to high internal  $[\text{H}^+]$ ). The influence of  $\text{HCO}_3^-$  channels on passive  $\text{HCO}_3^-$  fluxes from the chloroplast stroma into the thylakoid lumen was positively tested by means of a numerical model (Thoms et al., 2001) that takes into account the electrical potential difference across the thylakoid membrane. The electrochemical

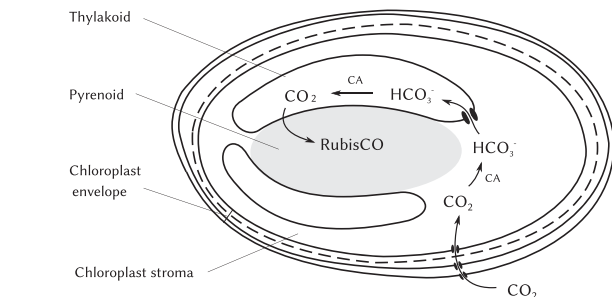
gradient established during illumination drives  $\text{H}^+$  out of the thylakoid, commonly through ATP synthetases. An import of  $\text{H}^+$  into the thylakoid is effected by linear or cyclic electron transport (cf. Raven et al., 2014). Inside the thylakoid lumen (pH ca. 5), CA activity is assumed. In green algae, a CA was found that is associated to the thylakoid membrane (Pronina and Semenenko, 1983). RubisCO is located inside the pyrenoid (cf. Fig. 3) (Vaughn et al., 1990; Borkhsenius et al., 1998). The pyrenoid of *E. huxleyi* is closely surrounded by thylakoid stacks composed of three thylakoid layers (Billard and Inouye, 2004). For the model set-up, we assume that thylakoid and pyrenoid are closely associated to each other and form a so-called thylakoid/pyrenoid complex. The  $\text{CO}_2$  permeability between thylakoid/pyrenoid lumen and the chloroplast stroma is set to a low value (ten-times lower than the value measured by Prasad et al. (1998) for pure liposomemembranes) which may be achieved by the characteristic arrangement of thylakoid stacks and pyrenoid, but also by a specialised composition of the thylakoid membranes. The permeability coefficients of thylakoid membranes have not yet been determined to the best of our knowledge. RubisCO is located also inside the *in silico* thylakoid/pyrenoid complex.

Aiming to re-enact the experiment of Rokitta and Rost (2012) (matrix approach with two different light and  $\text{CO}_2$  levels), the model is forced by four different external carbonate systems (Tab. 1), while setting the fixation rates of  $\text{CO}_2$  into POC (inside the thylakoid/pyrenoid complex) and  $\text{CO}_3^{2-}$  into PIC (inside the CV) to the POC and PIC production values measured by Rokitta and Rost (2012). In terms of POC production, an additional amount of  $\text{CO}_2$  is fixed that will be lost to respiration and thus be released into the cytosol. Respiration further provides the energy equivalents that are required for the  $\text{Ca}^{2+}/\text{HCO}_3^-/\text{H}^+$  transport across the CV membrane and therewith regulates PIC production rates.

## 2.1. Mathematical implementation

### 2.1.1. Carbonate system

The carbonate system ( $\text{CO}_2$ ,  $\text{HCO}_3^-$ ,  $\text{CO}_3^{2-}$ ,  $\text{H}^+$ , and  $\text{OH}^-$ ) (cf. Zeebe and Wolf-Gladrow, 2001) is resolved dynamically for all four compartments.



**Fig. 3.** Hypothetical chloroplast-based CCM for *Emiliania huxleyi*. Working hypothesis (similar to Pronina and Semenenko, 1992; Pronina and Borodin, 1993; Raven, 1997): Carbon dioxide enters the chloroplast from the cytosol via aquaporin-based diffusion. The high pH inside the chloroplast stroma (ca. 8) supports the carbonic anhydrase (CA)-catalysed conversion of  $\text{CO}_2$  to  $\text{HCO}_3^-$ . Bicarbonate diffuses into the thylakoid through channels. The low pH inside the thylakoid (ca. 5) supports CA-catalysed conversion of  $\text{HCO}_3^-$  to  $\text{CO}_2$ . The latter molecule diffuses into the pyrenoid where RubisCO is located.

**Table 1**  
Model boundary conditions. Concentrations of  $\text{CO}_2$  and  $\text{HCO}_3^-$  and pH in bulk seawater (temperature = 15 °C and salinity = 35), calculated from dissolved inorganic carbonate concentrations and pH measured by Rokitta and Rost (2012). The production rates of particulate organic and inorganic carbon ( $R_{\text{POC}}$  and  $R_{\text{PIC}}$ ) measured by Rokitta and Rost (2012) constitute sinks for inorganic carbon. In terms of POC production, RubisCO fixes an additional amount of  $\text{CO}_2$  that is released later-on into the cytosol via respiration ( $R_{\text{R}}$ ). LLLC, LLHC, HLLC, and HLHC determine the four model experiments, where LL stands for low light ( $50 \mu\text{mol photons} \cdot \text{m}^{-2} \cdot \text{s}^{-1}$ ), HL for high light ( $300 \mu\text{mol photons} \cdot \text{m}^{-2} \cdot \text{s}^{-1}$ ), LC for low and HC for high carbon, respectively.

Model experiment	$[\text{CO}_2]^{\text{SW}} (\text{mol} \cdot \text{m}^{-3})$	$[\text{HCO}_3^-]^{\text{SW}} (\text{mol} \cdot \text{m}^{-3})$	pH	$R_{\text{POC}} + R_{\text{R}} (\text{mol} \cdot \text{cell}^{-1} \cdot \text{h}^{-1})$	$R_{\text{PIC}} (\text{mol} \cdot \text{cell}^{-1} \cdot \text{h}^{-1})$
LLLC	$13 \times 10^{-3}$	1.9	8.1	$(35.4 + 5.59) \times 10^{-15}$	$49.3 \times 10^{-15}$
LLHC	$30 \times 10^{-3}$	2.2	7.8	$(65.6 + 1.64) \times 10^{-15}$	$13 \times 10^{-15}$
HLLC	$13 \times 10^{-3}$	1.9	8.1	$(82.8 + 24.8) \times 10^{-15}$	$67.3 \times 10^{-15}$
HLHC	$37 \times 10^{-3}$	2.1	7.7	$(92.1 + 6.45) \times 10^{-15}$	$56.7 \times 10^{-15}$





The corresponding net reaction rates are denoted according to the numbers of the rate constants and describe the reactions (8) to (12) from right to left:

$$R_1 = -k_{+1}[\text{CO}_2] + k_{-1}[\text{H}^+][\text{HCO}_3^-] \quad (13)$$

$$R_4 = -k_{+4}[\text{OH}^-][\text{CO}_2] + k_{-4}[\text{HCO}_3^-] \quad (14)$$

$$R_5^{\text{H}^+} = -k_{+5}^{\text{H}^+}[\text{H}^+][\text{CO}_3^{2-}] + k_{-5}^{\text{H}^+}[\text{HCO}_3^-] \quad (15)$$

$$R_5^{\text{OH}^-} = -k_{+5}^{\text{OH}^-}[\text{OH}^-][\text{HCO}_3^-] + k_{-5}^{\text{OH}^-}[\text{CO}_3^{2-}][\text{H}_2\text{O}] \quad (16)$$

$$R_6 = -k_{+6} + k_{-6}[\text{H}^+][\text{OH}^-] \quad (17)$$

where concentrations are given in  $\text{mol} \cdot \text{m}^{-3}$  and rates (R) in  $\text{mol} \cdot \text{m}^{-3} \cdot \text{h}^{-1}$ . For the chloroplast stroma and the thylakoid lumen, CA activity is assumed, i.e.  $R_1$  (eqn.(13)) is multiplied by  $10^4$  (cf. Supuran and Scozzafava, 2007).

### 2.1.2. Differential equations

The cell model consists of 24 coupled differential equations that describe the rates by which the concentrations of  $\text{CO}_2$ ,  $\text{HCO}_3^-$ ,  $\text{CO}_3^{2-}$ ,  $\text{H}^+$ ,  $\text{OH}^-$  (in all compartments),  $\text{Ca}^{2+}$  and PIC (inside the CV), POC (inside the thylakoid lumen), and energy equivalents (EE, inside the cytosol) change. In the following, the differential equations are listed according to their compartmental affiliation, where concentrations are given in  $\text{mol} \cdot \text{m}^{-3}$  and rates (R) in  $\text{mol} \cdot \text{m}^{-3} \cdot \text{h}^{-1}$ .

#### 2.1.2.1. Cytosol (CS).

$$\begin{pmatrix} d[\text{CO}_2]^{\text{CS}}/dt \\ d[\text{HCO}_3^-]^{\text{CS}}/dt \\ d[\text{CO}_3^{2-}]^{\text{CS}}/dt \\ d[\text{H}^+]^{\text{CS}}/dt \\ d[\text{OH}^-]^{\text{CS}}/dt \\ d[\text{EE}]^{\text{CS}}/dt \end{pmatrix} = \begin{pmatrix} \text{C system} \\ 0 \end{pmatrix} + \begin{pmatrix} R_{\text{CO}_2,\text{PM}}^{\text{CS}} - R_{\text{CO}_2,\text{CVm}}^{\text{CS}} - R_{\text{CO}_2,\text{CPm}}^{\text{CS}} + R_{\text{R}} \\ R_{\text{HCO}_3^-,\text{PM}}^{\text{CS}} - R_{\text{Ca}^{2+},\text{CVm}}^{\text{CS}} \\ 0 \\ R_{\text{Ca}^{2+},\text{CVm}}^{\text{CS}} - R_{\text{H}^+,\text{PM}}^{\text{CS}} - R_{\text{H}^+,\text{CPm}}^{\text{CS}} \\ 0 \\ 4.5 R_{\text{R}} - \frac{1}{2} R_{\text{Ca}^{2+},\text{CVm}}^{\text{CS}} \end{pmatrix} \quad (18)$$

“C system” stands for carbonate system which is described by the reaction rates given in equations (13)–(17).  $R_{\text{CO}_2,\text{PM}}$  stands for the  $\text{CO}_2$  -  $R_{\text{HCO}_3^-,\text{PM}}$  for the  $\text{HCO}_3^-$ , and  $R_{\text{H}^+,\text{PM}}$  for the  $\text{H}^+$  flux rate across the plasma membrane (1, 2, and 3 in Fig. 2).  $R_{\text{CO}_2,\text{CVm}}$  gives the  $\text{CO}_2$  flux rate and  $R_{\text{Ca}^{2+},\text{CVm}}$  the flux rate of  $\text{Ca}^{2+}$ ,  $\text{HCO}_3^-$ , and  $\text{H}^+$  across the CV membrane (4 and 5 in Fig. 2).  $R_{\text{CO}_2,\text{CPm}}$  and  $R_{\text{H}^+,\text{CPm}}$  denote the  $\text{CO}_2$ - and  $\text{H}^+$  flux rate across the chloroplast envelope, respectively (6 and 7 in Fig. 2).  $R_{\text{R}}$  is the respiration rate that releases  $\text{CO}_2$  into the cytosol while generating 4.5 energy equivalents (EE) per released  $\text{CO}_2$  (13 in Fig. 2, value approximated for ATP production).  $R_{\text{R}}$  is chosen so that the generated energy equivalents which drive the import of  $\text{Ca}^{2+}$  and  $\text{HCO}_3^-$  for calcite precipitation into the CV suffice to reach the PIC production rates measured by Rokitta and Rost (2012).

#### 2.1.2.2. Coccolith vesicle (CV)

$$\begin{pmatrix} d[\text{CO}_2]^{\text{CV}}/dt \\ d[\text{HCO}_3^-]^{\text{CV}}/dt \\ d[\text{CO}_3^{2-}]^{\text{CV}}/dt \\ d[\text{H}^+]^{\text{CV}}/dt \\ d[\text{OH}^-]^{\text{CV}}/dt \\ d[\text{Ca}^{2+}]^{\text{CV}}/dt \\ d[\text{PIC}]^{\text{CV}}/dt \end{pmatrix} = \begin{pmatrix} \text{C system} \\ 0 \\ 0 \end{pmatrix} + \begin{pmatrix} R_{\text{CO}_2,\text{CVm}}^{\text{CV}} \\ R_{\text{Ca}^{2+},\text{CVm}}^{\text{CV}} \\ -R_{\text{P}} \\ -R_{\text{Ca}^{2+},\text{CVm}}^{\text{CV}} \\ 0 \\ R_{\text{Ca}^{2+},\text{CVm}}^{\text{CV}} - R_{\text{P}} \\ R_{\text{P}} \end{pmatrix} \quad (19)$$

where  $R_{\text{P}}$  stands for the precipitation rate (11 in Fig. 2).

#### 2.1.2.3. Chloroplast stroma (CPstr)

$$\begin{pmatrix} d[\text{CO}_2]^{\text{CPstr}}/dt \\ d[\text{HCO}_3^-]^{\text{CPstr}}/dt \\ d[\text{CO}_3^{2-}]^{\text{CPstr}}/dt \\ d[\text{H}^+]^{\text{CPstr}}/dt \\ d[\text{OH}^-]^{\text{CPstr}}/dt \end{pmatrix} = \begin{pmatrix} \text{C system} \end{pmatrix} + \begin{pmatrix} R_{\text{CO}_2,\text{CPm}}^{\text{CPstr}} - R_{\text{CO}_2,\text{thyl}}^{\text{CPstr}} \\ -R_{\text{HCO}_3^-,\text{thyl}}^{\text{CPstr}} \\ 0 \\ R_{\text{H}^+,\text{CPm}}^{\text{CPstr}} - R_{\text{H}^+,\text{thyl}}^{\text{CPstr}} \\ 0 \end{pmatrix} \quad (20)$$

where the “C system” includes CA activity (cf. Section 2.1.1).  $R_{\text{CO}_2,\text{thyl}}$ ,  $R_{\text{HCO}_3^-,\text{thyl}}$  and  $R_{\text{H}^+,\text{thyl}}$  indicate the fluxes of  $\text{CO}_2$ ,  $\text{HCO}_3^-$  and  $\text{H}^+$  between chloroplast stroma and thylakoid lumen (8, 9, and 10 in Fig. 2).

#### 2.1.2.4. Thylakoid lumen (thyl)

$$\begin{pmatrix} d[\text{CO}_2]^{\text{thyl}}/dt \\ d[\text{HCO}_3^-]^{\text{thyl}}/dt \\ d[\text{CO}_3^{2-}]^{\text{thyl}}/dt \\ d[\text{H}^+]^{\text{thyl}}/dt \\ d[\text{OH}^-]^{\text{thyl}}/dt \\ d[\text{POC}]^{\text{thyl}}/dt \end{pmatrix} = \begin{pmatrix} \text{C system} \\ 0 \end{pmatrix} + \begin{pmatrix} R_{\text{CO}_2}^{\text{thyl}} - R_{\text{CO}_2}^{\text{RubisCO}} \\ 0 \\ 0 \\ R_{\text{H}^+}^{\text{thyl}} \\ 0 \\ R_{\text{CO}_2}^{\text{RubisCO}} \end{pmatrix} \quad (21)$$

where the “C system” also includes CA activity  $R_{\text{CO}_2}^{\text{RubisCO}}$  gives the  $\text{CO}_2$  fixation rate by RubisCO (12 in Fig. 2) which is set to the POC production rate measured by Rokitta and Rost (2012) plus the respiration rate  $R_{\text{R}}$ .

### 2.1.3. Flux rates

For the calculation of some fluxes, we use smooth cut-off functions ( $\text{fun}_X$ ) that rely on the following structure:

$$\text{fun}_X = Y_i \cdot \tanh(X - X_i) + Y_{ii} \quad (22)$$

where  $X$  denotes the variable value such as for instance the cytosolic pH in equation (25),  $X_i$  indicates the  $X$  value, where the strongest change in  $\text{fun}_X$  occurs.  $Y_i$  and  $Y_{ii}$  define the size range of  $\text{fun}_X$ . Values for  $X_i$ ,  $Y_i$ , and  $Y_{ii}$  are given in Table 2.

In the following, flux rate calculations are listed in the order of the numbers allocated in Figure 2. Concentrations are given in  $\text{mol} \cdot \text{m}^{-3}$ , rates in  $\text{mol} \cdot \text{m}^{-3} \cdot \text{h}^{-1}$ , permeability coefficients ( $\gamma$ ) in  $\text{m} \cdot \text{h}^{-1}$ , volumes (V) in  $\text{m}^3$ , and surface areas (A) in  $\text{m}^2$ . External concentrations of  $\text{CO}_2$ ,  $\text{HCO}_3^-$  and  $\text{H}^+$  are listed in Table 1. Volumes and surface areas of the four compartments are listed in Table 3. Physiological parameter values and a description of the parameters are given in Table 4.

#### (1) $\text{CO}_2$ flux across plasma membrane

$R_{\text{CO}_2,\text{PM}}^{\text{CS}}$  gives the rate by which the  $\text{CO}_2$  concentration inside the cytosol (CS) changes due to  $\text{CO}_2$  diffusion across the plasma membrane (PM).

$$R_{\text{CO}_2,\text{PM}}^{\text{CS}} = \gamma_{\text{CO}_2,\text{PM}} \cdot \frac{A^{\text{CS}}}{V^{\text{CS}}} \cdot ([\text{CO}_2]^{\text{SW}} - [\text{CO}_2]^{\text{CS}}) \quad (23)$$

#### (2) $\text{HCO}_3^-$ flux across plasma membrane

$$R_{\text{HCO}_3^-,\text{PM}}^{\text{CS}} = \gamma_{\text{HCO}_3^-,\text{PM}} \cdot \frac{A^{\text{CS}}}{V^{\text{CS}}} \cdot ([\text{HCO}_3^-]^{\text{SW}} - [\text{HCO}_3^-]^{\text{CS}}) \quad (24)$$

#### (3.) Proton flux across plasma membrane

Protons enter or leave the cell in order to keep cytosolic pH close to 7. The rate by which the  $\text{H}^+$  concentration inside the cytosol changes is:

$$R_{\text{H}^+,\text{PM}}^{\text{CS}} = \text{fun}_{\text{H}^+} \cdot R_{\text{max},\text{PM}}^{\text{H}^+} \quad (25)$$

(4.) CO<sub>2</sub> flux across coccolith vesicle membrane

$$R_{\text{CO}_2, \text{CVm}}^{\text{CV}} = \gamma_{\text{CO}_2, \text{CVm}} \cdot \frac{A^{\text{CV}}}{V^{\text{CV}}} \cdot ([\text{CO}_2]^{\text{CS}} - [\text{CO}_2]^{\text{CV}}) \quad (26)$$

The corresponding rate for the cytosol reads:

$$R_{\text{CO}_2, \text{CVm}}^{\text{CS}} = R_{\text{CO}_2, \text{CVm}}^{\text{CV}} \cdot \frac{V^{\text{CV}}}{V^{\text{CS}}} \quad (27)$$

**Table 2**

Parameter values for the cut-off functions. The basic structure of the functions is given by eqn. (22). The concentrations of Ca<sup>2+</sup>, energy equivalents (EE), and HCO<sub>3</sub><sup>-</sup> are given in mol · m<sup>-3</sup>. CPstr - chloroplast stroma, CS - cytosol, CV - coccolith vesicle, thyl - thylakoid lumen, column "eqn." hints at the equations, in which the cut-off functions are used.

Function	X	X <sub>i</sub>	Y <sub>i</sub>	Y <sub>ii</sub>	eqn.
fun <sub>Ca<sup>2+</sup></sub>	-100 · [Ca <sup>2+</sup> ] <sup>CV</sup>	-10 <sup>3</sup>	0.5	0.5	(28)
fun <sub>EE</sub>	100 · [EE] <sup>CS</sup>	2 × 10 <sup>4</sup>	0.5	0.5	(28)
fun <sub>H<sup>+</sup></sub>	pH <sup>CS</sup>	7	-1	0	(25)
fun <sub>H<sub>2</sub><sup>+</sup></sub>	pH <sup>CPstr</sup>	8	-1	0	(32)
fun <sub>H<sub>3</sub><sup>+</sup></sub>	pH <sup>thyl</sup>	5	-1	0	(38)
fun <sub>HCO<sub>3</sub><sup>-</sup></sub>	10 · [HCO <sub>3</sub> <sup>-</sup> ] <sup>CS</sup>	0.3	0.5	0.5	(28)

**Table 3**

Morphological parameter values used for the cell model. CP - chloroplast, CPstr - chloroplast stroma, CS - cytosol, CV - coccolith vesicle, PM - plasma membrane, thyl - thylakoid lumen, thyl - thylakoid/pyrenoid complex.

Parameter (unit)	Description	value
V <sup>CS</sup> (m <sup>3</sup> )	Volume CS	16.1 × 10 <sup>-18</sup>
A <sup>CS</sup> (m <sup>2</sup> )	PM amount	78.5 × 10 <sup>-12</sup>
V <sup>CV</sup> (m <sup>3</sup> )	Volume CV	1.6 × 10 <sup>-18</sup>
A <sup>CV</sup> (m <sup>2</sup> )	Membrane surrounding CV	16.8 × 10 <sup>-12</sup>
V <sup>CP</sup> (m <sup>3</sup> )	Volume of (bulk) CP	28.3 × 10 <sup>-18</sup>
V <sup>CPstr</sup> (m <sup>3</sup> )	Volume of CPstr	22.6 × 10 <sup>-18</sup>
A <sup>CP</sup> (m <sup>2</sup> )	Membrane surrounding CP	64.1 × 10 <sup>-12</sup>
V <sup>thyl</sup> (m <sup>3</sup> )	Volume of thylakoid	5.7 × 10 <sup>-18</sup>
A <sup>thyl</sup> (m <sup>2</sup> )	Membrane surrounding thylakoid	37 × 10 <sup>-12</sup>

**Table 4**

Physiological parameter values for the model. AQP - aquaporin, CPm - chloroplast envelope, CPstr - chloroplast stroma, CV - coccolith vesicle, CVm - CV membrane, POC - particulate organic carbon, PON - particulate organic nitrogen, PM - plasma membrane. LLLC, LLHC, HLLC, and HLHC denote the four model experiments where LL stands for low light, HL for high light, LC for low and HC for high carbon.

Parameter (unit)	Description	LLLC	LLHC	HLLC	HLHC
γ <sub>CO<sub>2</sub>,CPm</sub> (m · h <sup>-1</sup> )	Permeability coefficient of CPm for CO <sub>2</sub>	0.1 Uehlein et al. (2008)			
γ <sub>CO<sub>2</sub>,CVm</sub> (m · h <sup>-1</sup> )	Permeability coefficient of CVm for CO <sub>2</sub> (no AQPs)	0.02 Prasad et al. (1998)			
γ <sub>CO<sub>2</sub>,PM</sub> (m · h <sup>-1</sup> )	AQP-based permeability coefficient of PM for CO <sub>2</sub>	0.31 Uehlein et al. (2008)			
γ <sub>CO<sub>2</sub>,thyl</sub> (m · h <sup>-1</sup> )	Permeability coefficient of thylakoid/pyrenoid complex for CO <sub>2</sub>	0.002			
γ <sub>HCO<sub>3</sub><sup>-</sup>,PM</sub> (m · h <sup>-1</sup> )	Permeability coefficient of PM for HCO <sub>3</sub> <sup>-</sup>	0.0016			
γ <sub>HCO<sub>3</sub><sup>-</sup>,thyl</sub> (m · h <sup>-1</sup> )	Permeability coefficient of thylakoid/pyrenoid complex for CO <sub>2</sub>	0.02			
K <sub>m,Ca<sup>2+</sup>/HCO<sub>3</sub><sup>-</sup>/H<sup>+</sup></sub> (mol · m <sup>-3</sup> )	Half saturation constant of Ca <sup>2+</sup> /HCO <sub>3</sub> <sup>-</sup> /H <sup>+</sup> transporter towards cytosolic [Ca <sup>2+</sup> ]	10 <sup>-3</sup>			
k <sub>f</sub> (mol · m <sup>-3</sup> · h <sup>-1</sup> )	Calcite precipitation related rate constant	110 × 10 <sup>-9</sup> Zuddas and Mucci (1994)			
n	Calcite precipitation related parameter	2.35 Zuddas and Mucci (1994)			
R <sub>max,CPm</sub> <sup>H<sup>+</sup></sup> (mol · m <sup>-3</sup> · h <sup>-1</sup> )	Maximum rate by which [H <sup>+</sup> ] in CPstr changes due to H <sup>+</sup> fluxes across CPm	1.94 × 10 <sup>8</sup>			
R <sub>max,PM</sub> <sup>H<sup>+</sup></sup> (mol · m <sup>-3</sup> · h <sup>-1</sup> )	Maximum rate by which [H <sup>+</sup> ] in CS changes due to H <sup>+</sup> fluxes across PM	1.7 × 10 <sup>6</sup>			
R <sub>max,thyl</sub> <sup>H<sup>+</sup></sup> (mol · m <sup>-3</sup> · h <sup>-1</sup> )	Maximum rate by which [H <sup>+</sup> ] in thylakoid lumen changes	3.95 × 10 <sup>8</sup>			
R <sub>POC</sub> (mol · cell <sup>-1</sup> · h <sup>-1</sup> )	Net POC production rate Rokitta and Rost (2012)	35.4 × 10 <sup>-15</sup>	65.5 × 10 <sup>-15</sup>	82.8 × 10 <sup>-15</sup>	92.1 × 10 <sup>-15</sup>
R <sub>R</sub> (mol · cell <sup>-1</sup> · h <sup>-1</sup> )	Respiration rate	5.59 × 10 <sup>-15</sup>	1.64 × 10 <sup>-15</sup>	24.8 × 10 <sup>-15</sup>	6.45 × 10 <sup>-15</sup>
V <sub>max,Ca<sup>2+</sup>/HCO<sub>3</sub><sup>-</sup>/H<sup>+</sup></sub> (mol · h <sup>-1</sup> )	Maximal transport velocity of Ca <sup>2+</sup> /HCO <sub>3</sub> <sup>-</sup> /H <sup>+</sup> transporter	0.72 × 10 <sup>-12</sup>			

R<sub>CO<sub>2</sub>,CVm</sub><sup>CS</sup> is subtracted from the cytosolic CO<sub>2</sub> pool, when R<sub>CO<sub>2</sub>,CVm</sub><sup>CV</sup> is added to the CV-internal CO<sub>2</sub> pool (cf. eqns. (18) and (19)).

- (5.) Ca<sup>2+</sup>/HCO<sub>3</sub><sup>-</sup>/H<sup>+</sup> fluxes across coccolith vesicle membrane  
Calcium ions can be accumulated within the CV up to a CV-internal concentration of 10 mol · m<sup>-3</sup> (fun<sub>Ca<sup>2+</sup></sub>) as previously presumed in Holtz et al. (2013). Furthermore, the Ca<sup>2+</sup>/HCO<sub>3</sub><sup>-</sup>/H<sup>+</sup> transport (stoichiometry: 1:1:1, cf. (eqns. (18) and 19)) is made dependent on the availability of cytosolic HCO<sub>3</sub><sup>-</sup> and energy equivalents (fun<sub>HCO<sub>3</sub><sup>-</sup></sub>, fun<sub>EE</sub>). The rate by which the Ca<sup>2+</sup> concentration inside the CV changes is:

$$R_{\text{Ca}^{2+}, \text{CVm}}^{\text{CV}} = \text{fun}_{\text{Ca}^{2+}} \cdot \text{fun}_{\text{HCO}_3^-} \cdot \text{fun}_{\text{EE}} \cdot \frac{1}{V^{\text{CV}}} \cdot \frac{V_{\text{max}, \text{Ca}^{2+}/\text{HCO}_3^-/\text{H}^+} \cdot [\text{Ca}^{2+}]^{\text{CS}}}{K_{\text{m}, \text{Ca}^{2+}/\text{HCO}_3^-/\text{H}^+} + [\text{Ca}^{2+}]^{\text{CS}}} \quad (28)$$

For the cytosol, the rate is:

$$R_{\text{Ca}^{2+}, \text{CVm}}^{\text{CS}} = R_{\text{Ca}^{2+}, \text{CVm}}^{\text{CV}} \cdot \frac{V^{\text{CV}}}{V^{\text{CS}}} \quad (29)$$

The Ca<sup>2+</sup> concentration inside the cytosol ([Ca<sup>2+</sup>]<sup>CS</sup>) is set to a fixed value, i.e. 10<sup>-4</sup> mol · m<sup>-3</sup> (Brownlee et al., 1995).

- (6.) CO<sub>2</sub> flux across chloroplast envelope

$$R_{\text{CO}_2, \text{CPm}}^{\text{CPstr}} = \gamma_{\text{CO}_2, \text{CPm}} \cdot \frac{A^{\text{CP}}}{V^{\text{CPstr}}} \cdot ([\text{CO}_2]^{\text{CS}} - [\text{CO}_2]^{\text{CPstr}}) \quad (30)$$

The corresponding rate for the cytosol reads:

$$R_{\text{CO}_2, \text{CPm}}^{\text{CS}} = R_{\text{CO}_2, \text{CPm}}^{\text{CPstr}} \cdot \frac{V^{\text{CPstr}}}{V^{\text{CS}}} \quad (31)$$

- (7.) Proton flux across chloroplast envelope

Protons enter or leave the chloroplast stroma in order to keep the pH close to 8 (fun<sub>H<sup>+</sup></sub><sub>2</sub>).

$$R_{\text{H}^+, \text{CPm}}^{\text{CPstr}} = \text{fun}_{\text{H}_2^+} \cdot R_{\text{max}, \text{CPm}}^{\text{H}^+} \quad (32)$$

The corresponding rate for the cytosol reads:

$$R_{\text{H}^+, \text{CPm}}^{\text{CS}} = R_{\text{H}^+, \text{CPm}}^{\text{CPstr}} \cdot \frac{V^{\text{CPstr}}}{V^{\text{CS}}} \quad (33)$$

- (8.)
- $\text{HCO}_3^-$
- flux between chloroplast stroma and thylakoid lumen

$$R_{\text{HCO}_3^-}^{\text{thyl}} = \gamma_{\text{HCO}_3^-, \text{thyl}} \cdot \frac{A^{\text{thyl}}}{V^{\text{thyl}}} \cdot ([\text{HCO}_3^-]^{\text{CPstr}} - [\text{HCO}_3^-]^{\text{thyl}}) \quad (34)$$

The corresponding rate for the chloroplast stroma reads:

$$R_{\text{HCO}_3^-}^{\text{CPstr}} = R_{\text{HCO}_3^-}^{\text{thyl}} \cdot \frac{V^{\text{thyl}}}{V^{\text{CPstr}}} \quad (35)$$

- (9.)
- $\text{CO}_2$
- flux between chloroplast stroma and thylakoid lumen

$$R_{\text{CO}_2}^{\text{thyl}} = \gamma_{\text{CO}_2, \text{thyl}} \cdot \frac{A^{\text{thyl}}}{V^{\text{thyl}}} \cdot ([\text{CO}_2]^{\text{CPstr}} - [\text{CO}_2]^{\text{thyl}}) \quad (36)$$

The corresponding rate for the chloroplast stroma reads:

$$R_{\text{CO}_2}^{\text{CPstr}} = R_{\text{CO}_2}^{\text{thyl}} \cdot \frac{V^{\text{thyl}}}{V^{\text{CPstr}}} \quad (37)$$

- (10.) Proton flux between chloroplast stroma and thylakoid lumen
- 
- Protons enter or leave the thylakoid lumen in order to keep the pH of the thylakoid lumen close to 5 (
- $\text{fun}_{\text{H}^+}$
- ).

$$R_{\text{H}^+}^{\text{thyl}} = \text{fun}_{\text{H}^+} \cdot R_{\text{max}, \text{thyl}}^{\text{H}^+} \quad (38)$$

The corresponding rate for the chloroplast stroma reads:

$$R_{\text{H}^+}^{\text{CPstr}} = R_{\text{H}^+}^{\text{thyl}} \cdot \frac{V^{\text{thyl}}}{V^{\text{CPstr}}} \quad (39)$$

- (11.) Calcite precipitation inside coccolith vesicle
- 
- Calcite is precipitated from
- $\text{Ca}^{2+}$
- and
- $\text{CO}_3^{2-}$
- :



The corresponding precipitation rate (after Zuddas and Mucci, 1994) is:

$$R_p = \begin{cases} k_f(\Omega - 1)^n & \text{for } \Omega > 1 \\ 0 & \text{for } \Omega \leq 1 \end{cases} \quad (41)$$

where  $\Omega$  is the calcite saturation product that is dependent on the prevailing concentrations of  $\text{Ca}^{2+}$  and  $\text{CO}_3^{2-}$ . The parameter  $n$  is given in Zuddas and Mucci (1994), and  $k_f$  can be calculated (cf. Zuddas and Mucci, 1994; Holtz et al., 2013).

- (12.)
- $\text{CO}_2$
- fixation inside thylakoid/pyrenoid complex
- 
- The
- $\text{CO}_2$
- fixation rate of RubisCO (
- $R_{\text{CO}_2}^{\text{RubisCO}}$
- ) is calculated as follows:

$$R_{\text{CO}_2}^{\text{RubisCO}} = \frac{1}{V^{\text{thyl}}} (R_{\text{POC}} + R_R) \quad (42)$$

- (13.) Respiration inside cytosol
- 
- Respiration releases
- $\text{CO}_2$
- into the cytosol with the following rate.

$$R_R = \frac{1}{V^{\text{CS}}} R_{\text{R}} \quad (43)$$

### 3. Results

The model was run into steady state, i.e. until all state variables remained constant over time. The values calculated for experiments LLLC, LLHC, and HLHC are listed in Table 5. The  $\text{CO}_2$  and  $\text{HCO}_3^-$  concentration gradients that are obtained direct from seawater across the cytoplasm into the chloroplast stroma. While  $\text{CO}_2$  can diffuse across the plasma membrane and the chloroplast envelope,  $\text{HCO}_3^-$  cannot cross the chloroplast envelope in our

**Table 5**

Calculated steady state values inside the four model compartments. Concentrations are given in  $\text{mol} \cdot \text{m}^{-3}$ ,  $\Omega$  is dimensionless. LLLC, LLHC, and HLHC denote the three model experiments that can be explained by the fluxes proposed in Fig. 2 where LL stands for low light, HL for high light, LC for low and HC for high carbon.

	LLC	LLHC	HLHC
<b>Cytoplasm</b>			
pH	7.00	7.00	7.00
$\text{CO}_2$	$11.7 \times 10^{-3}$	$27.8 \times 10^{-3}$	$33.4 \times 10^{-3}$
$\text{HCO}_3^-$	1.44	2.03	1.61
$\text{CO}_3^{2-}$	$10.2 \times 10^{-3}$	$14.4 \times 10^{-3}$	$11.4 \times 10^{-3}$
<b>Coccolith vesicle</b>			
pH	8.62	8.32	8.34
$\text{CO}_2$	$11.7 \times 10^{-3}$	$27.8 \times 10^{-3}$	$33.4 \times 10^{-3}$
$\text{HCO}_3^-$	5.12	6.22	7.74
$\text{CO}_3^{2-}$	1.46	0.915	1.16
$\text{Ca}^{2+}$	3.10	3.10	4.11
$\Omega$	12.1	7.59	12.8
<b>Chloroplast stroma</b>			
pH	8.00	8.00	8.00
$\text{CO}_2$	$5.5 \times 10^{-3}$	$17.6 \times 10^{-3}$	$18.5 \times 10^{-3}$
$\text{HCO}_3^-$	$77.3 \times 10^{-3}$	0.323	0.307
$\text{CO}_3^{2-}$	$5.49 \times 10^{-3}$	$23.0 \times 10^{-3}$	$21.8 \times 10^{-3}$
<b>Thylakoid/pyrenoid complex</b>			
pH	5.00	5.00	5.00
$\text{CO}_2$	0.101	1.10	0.818
$\text{HCO}_3^-$	$12.7 \times 10^{-3}$	0.124	$94.4 \times 10^{-3}$
$\text{CO}_3^{2-}$	$0.9 \times 10^{-6}$	$8.9 \times 10^{-6}$	$6.8 \times 10^{-6}$

model. Inside the thylakoid/pyrenoid complex, the  $\text{HCO}_3^-$  concentration is lower than inside the chloroplast stroma, while  $\text{CO}_2$  is strongly accumulated. The passage of  $\text{HCO}_3^-$  across the thylakoid/pyrenoid complex-confining membrane is facilitated via channels;  $\text{CO}_2$  diffusion, in contrast, is low.

Under high light intensities and low external  $\text{CO}_2$  concentrations (HLLC), some of the state variable concentrations inside the chloroplast become negative when using the parameter set given in Table 4 and the model does not reach steady state until the end of the model run (three days in total). It follows that cellular carbon fluxes under HLLC conditions cannot be explained by the fluxes proposed in Figure 2. A hypothesis will be developed in the discussion section (Section 4.2) that considers the variation of cellular carbon fluxes under different light and  $\text{CO}_2$  conditions and hence includes fluxes under HLLC conditions.

### 4. Discussion

The discussion is presented in the sequence of the three questions posed at the end of the introduction.

#### 4.1. Can $\text{CO}_2$ and $\text{HCO}_3^-$ pass the cytosol without being inter-converted?

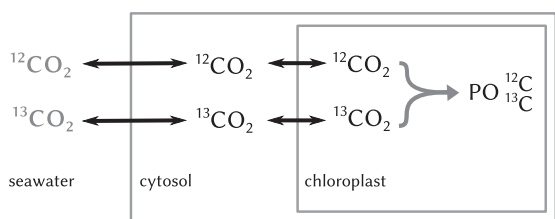
The bulk of experiments conducted until today gives evidence that POC production in *E. huxleyi* is basically fed by external  $\text{CO}_2$  and external  $\text{HCO}_3^-$  and that PIC production is fed by external  $\text{HCO}_3^-$  only. Thus, we suggested in the beginning that  $\text{CO}_2$  is rarely converted to  $\text{HCO}_3^-$  within the cytosol, because PIC production seems to be fed by external  $\text{HCO}_3^-$  only.

Compared to influx rates of  $\text{CO}_2$  and  $\text{HCO}_3^-$ , calculated inter-conversion rates between cytosolic  $\text{CO}_2$  and  $\text{HCO}_3^-$  are low (Tab. 6). Due to the higher concentrations of  $\text{HCO}_3^-$  compared to  $\text{CO}_2$ , conversion rates of  $\text{HCO}_3^-$  to  $\text{CO}_2$  are higher than those of the inverse reaction. As mentioned above, the carbon demand of photosynthesis in experiment HLLC cannot be covered by means of the carbon fluxes proposed in Figure 2. We find that neither

**Table 6**

Inter-conversion rates between cytosolic  $\text{CO}_2$  and  $\text{HCO}_3^-$  as percentage of the influx of  $\text{CO}_2$  and  $\text{HCO}_3^-$ , respectively. Rel. conv. of  $\text{CO}_2$  to  $\text{HCO}_3^-$  (%) - conversion rate of cytosolic  $\text{CO}_2$  into  $\text{HCO}_3^-$  as percentage of cellular  $\text{CO}_2$  influx; Rel. conv. of  $\text{HCO}_3^-$  to  $\text{CO}_2$  (%) - conversion rate of cytosolic  $\text{HCO}_3^-$  into  $\text{CO}_2$  as percentage of  $\text{HCO}_3^-$  cellular influx. LLLC, LLHC, and HLHC denote the three model experiments that lead to reasonable results where LL stands for low light, HL for high light, LC for low and HC for high carbon.

	LLLC	LLHC	HLHC
Rel. conv. of $\text{CO}_2$ to $\text{HCO}_3^-$ (%)	0.03	0.04	0.03
Rel. conv. of $\text{HCO}_3^-$ to $\text{CO}_2$ (%)	0.24	1.16	0.24



**Fig. 4.** Stable carbon isotope model.  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  diffuse independently of each other across plasma membrane and chloroplast envelope. The chloroplast is surrounded by three membranes which is not indicated in the figure. Inside the chloroplast,  $\text{CO}_2$  is fixed into particulate organic carbon (POC) with the  $\text{PO } ^{12}\text{C}$  and  $\text{PO } ^{13}\text{C}$  fixation rates measured by Rost et al. (2002). The concentrations of external  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  were measured by Rost et al. (2002) also.

CA-activity nor strong  $\text{HCO}_3^-$  accumulation ( $> 150 \text{ mol} \cdot \text{m}^{-3}$ ) inside the cytosol are able to provide photosynthesis with enough  $\text{CO}_2$  stemming from cytosolic  $\text{HCO}_3^-$  conversion. Assuming both at a time, CA-activity and  $\text{HCO}_3^-$  accumulation inside the cytosol ( $\sim 150 \text{ mol} \cdot \text{m}^{-3}$ ), carbon provision from cytosolic  $\text{HCO}_3^-$  conversion suffices to cover the demand of photosynthesis in experiment HLLC. Under these conditions, however, around 99% of the imported  $\text{HCO}_3^-$  is lost from the cell via  $\text{CO}_2$  diffusion which would imply very high energy demands. An alternative, more energy-efficient carbon pathway will be proposed in Section 4.2 (cf. Fig. 5).

#### 4.2. Can enough external $\text{CO}_2$ be taken up by diffusion to satisfy the carbon requirement of POC production?

Diffusive  $\text{CO}_2$  influx relies on a  $\text{CO}_2$  concentration gradient from the external medium into the cytoplasm. In our model, such a gradient is established via an energy-efficient, chloroplast-based CCM that relies on internal pH gradients, CA-activity inside the chloroplast stroma and the thylakoid lumen,  $\text{CO}_2$  fixation by RubisCO,  $\text{HCO}_3^-$  channels inside the thylakoid membrane, and differences in  $\text{CO}_2$  permeabilities between the various membranes (for a review on energy-costs of CCMs see Raven et al., 2014).

Carbon dioxide permeabilities of plasma membrane and chloroplast envelope play a decisive role for the diffusive influx of  $\text{CO}_2$ . Our calculations are based on the  $\text{CO}_2$  permeability coefficient determined for the plasma membrane of *N. tabacum* that is known to include aquaporins ( $0.31 \text{ m} \cdot \text{h}^{-1}$ ) (Uehlein et al., 2008). On basis of the (lower)  $\text{CO}_2$  permeability coefficients determined for *Chlamydomonas reinhardtii* ( $0.03 - 0.06 \text{ m} \cdot \text{h}^{-1}$ ) (Sültemeyer and Rinast, 1996) and for artificial membranes that contain aquaporins from human erythrocytes ( $0.07 \text{ m} \cdot \text{h}^{-1}$ ) (Prasad et al., 1998), calculated fluxes of external  $\text{CO}_2$  towards RubisCO are lower than the POC production rates determined by Rokitta and Rost (2012) in all four model simulations. It follows that the proposed CCM depends on comparably high  $\text{CO}_2$  permeabilities of the plasma membrane and the chloroplast envelope. These permeabilities have not yet been determined for *E. huxleyi*. The  $\text{CO}_2$  permeability coefficients determined by Hopkinson et al. (2011) for the plasma membranes of different diatom species, however, are in the range

from  $0.54$  to  $2 \text{ m} \cdot \text{h}^{-1}$  and thus exceed the value used in our model. Nevertheless, using even the highest permeability value determined by Hopkinson et al. (2011) does not lead to sufficiently high carbon fluxes in model experiment HLLC.

In contrast to our assumption that  $\text{CO}_2$  enters the cell following its concentration gradient, some authors (e.g. Schulz et al., 2007) assume that  $\text{CO}_2$  leaves the cell of *E. huxleyi* following a concentration gradient that is established via an active import of  $\text{CO}_2$ . This  $\text{CO}_2$  efflux is thought to support the export of  $^{13}\text{CO}_2$  that accumulates inside the cell due to RubisCO's fractionation against  $^{13}\text{CO}_2$ . A simple two-compartment model is established (Fig. 4) to test whether a net efflux of  $^{13}\text{CO}_2$  is necessary for the cell to dispose of  $^{13}\text{C}$  that accumulates around RubisCO.

The model takes into account diffusive fluxes of  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  between medium/seawater, cytosol, and "bulk chloroplast" (i.e. chloroplast is not divided into chloroplast stroma and thylakoid/pyrenoid complex as in the other model). Inside the bulk chloroplast,  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  are fixed into POC with the rates measured by Rost et al. (2002). Rost et al. (2002) measured the fixation rates of  $^{12}\text{C}$  and  $^{13}\text{C}$  into POC in dependence on different external  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  concentrations and light conditions. The mathematical implementation of our model is given in the Appendix.

The model shows that *E. huxleyi* does not rely on a net efflux of  $^{13}\text{CO}_2$  in order to dispose of the  $^{13}\text{CO}_2$  that may accumulate inside the cell due to RubisCO's fractionation against  $^{13}\text{CO}_2$ . Calculated concentration gradients for  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  direct from the external medium/seawater into the bulk chloroplast (Tab. 7) and the resulting  $\text{CO}_2$  influx has a lower  $^{13}\text{R}_{\text{CO}_2}$  ( $[^{13}\text{CO}_2]:[^{12}\text{CO}_2]$ ) than the external medium ( $\Delta^{13}\text{R}_{\text{CO}_2,\text{infl}}$  in Tab. 7), because the ratio between the  $^{13}\text{CO}_2$  fixation rate of RubisCO and the external  $^{13}\text{CO}_2$  concentration is lower than the same ratio for  $^{12}\text{CO}_2$ .

For 29 out of 35 model experiments, the model shows that diffusive  $\text{CO}_2$  uptake can explain the POC production rates measured by Rost et al. (2002) (Tab. 7). The remaining six data points (italic numbers in Tab. 7) represent experiments that were conducted under low external  $\text{CO}_2$  concentrations and high light intensities which is well in line with the findings presented in Section 4.2 for the cell model. We hence argue that whether diffusive  $\text{CO}_2$  uptake can cover the carbon demand of RubisCO or not depends on the light intensity on the one hand and on the availability of external  $\text{CO}_2$  on the other. In case the carbon demand of RubisCO cannot be covered by  $\text{CO}_2$  diffusion only, we propose an additional uptake of  $\text{HCO}_3^-$  (cf. hypothesis presented in Fig. 5) which is in line with the findings of Rost et al. (2002); Bach et al. (2013), and Kottmeier et al.: The stable carbon isotope composition of POC measured by Rost et al. (2002) under different light and  $\text{CO}_2$  levels can be explained by an additional uptake of  $\text{HCO}_3^-$  when  $\text{CO}_2$  availability is low. Bach et al. (2013) showed by means of transcriptome analyses that anion exchangers that are presumably involved in  $\text{HCO}_3^-$  uptake are upregulated under low external DIC concentrations. Kottmeier et al. measured  $\text{HCO}_3^-$  usage for POC production only under low external  $\text{CO}_2$  concentrations. Under high  $\text{CO}_2$  concentrations, POC production relies on  $\text{CO}_2$  uptake exclusively. As shown in Section 4.1, high conversion rates of cytosolic  $\text{HCO}_3^-$  to  $\text{CO}_2$  that could feed the chloroplast with  $\text{CO}_2$  are dependent on CA-activity and a very high  $\text{HCO}_3^-$  accumulation inside the cytosol which would then lead to very high  $\text{CO}_2$  efflux rates. An additional uptake of  $\text{HCO}_3^-$  into the cytosol and from there into the chloroplast would certainly constitute the more energy-efficient possibility to import carbon. According to the state variable concentrations in steady state calculated for the four model compartments (Tab. 5),  $\text{HCO}_3^-$  could even follow its concentration gradient into the chloroplast.

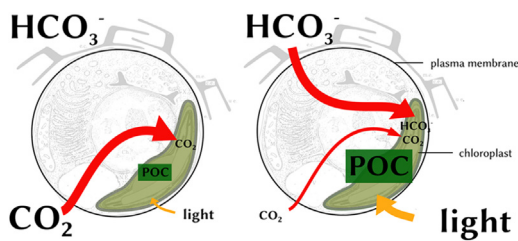
To sum up, depending on the  $\text{CO}_2$  permeabilities of plasma membrane and chloroplast envelope,  $\text{CO}_2$  for POC production can be taken up via diffusion under LLLC, LLHC, and HLHC, but not under HLLC conditions.



**Table 7**

Model outputs to stable carbon isotope model.  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  concentrations in the cytosol (CS, in steady state), and the bulk chloroplast (CP). All concentrations are given in  $\text{mmol} \cdot \text{m}^{-3}$ . Negative concentrations inside the chloroplast indicate that a diffusive influx of external  $\text{CO}_2$  cannot be the only carbon source for POC production. L:D - light: dark cycle (unit: h:h); PFD - photon flux density (unit:  $\mu\text{mol photons} \cdot \text{m}^{-2} \cdot \text{s}^{-1}$ ).  $\Delta^{13}\text{R}_{\text{CO}_2,\text{infl}}$  stands for:  $10^3 \cdot ([^{13}\text{CO}_2] : [^{12}\text{CO}_2])_{\text{ext. medium}} - ([^{13}\text{CO}_2] : [^{12}\text{CO}_2])_{\text{cell. inflx}}$ ; a positive value thus indicates that  $[^{13}\text{CO}_2] : [^{12}\text{CO}_2]$  of ext. medium exceeds  $[^{13}\text{CO}_2] : [^{12}\text{CO}_2]$  of cellular  $\text{CO}_2$  influx. Numbers set in italic indicate the six experiments that cannot be explained by the model.

L:D	PFD	CS		CP		$\Delta^{13}\text{R}_{\text{CO}_2,\text{infl}}$
		$^{12}\text{CO}_2$	$^{13}\text{CO}_2$	$^{12}\text{CO}_2$	$^{13}\text{CO}_2$	
24:0	15	32.6	0.363	29.7	0.331	0.166
		29.3	0.326	26.7	0.297	0.163
		18.3	0.203	15.6	0.174	0.161
		18.9	0.211	16.2	0.181	0.160
		12.3	0.137	9.7	0.109	0.174
		25.4	0.282	19.4	0.217	0.168
	30	21.6	0.241	16.4	0.184	0.175
		17.1	0.191	12.4	0.139	0.170
		15.3	0.171	10.3	0.115	0.173
		10.6	0.118	6.5	0.073	0.178
		25.4	0.284	12.8	0.146	0.185
		24.2	0.270	13.9	0.157	0.187
	80	16.6	0.186	6.9	0.079	0.195
		16.8	0.187	8.2	0.093	0.184
		9.4	0.105	1.6	0.019	0.178
		27.0	0.300	17.0	0.192	0.193
		19.9	0.222	11.1	0.126	0.187
		14.8	0.165	4.3	0.050	0.187
	150	14.0	0.157	4.6	0.053	0.191
		9.3	0.104	1.0	0.013	0.185
16:8	30	30.6	0.340	26.2	0.292	0.086
		22.4	0.249	18.2	0.203	0.074
		14.0	0.155	6.4	0.072	0.083
		9.2	0.103	3.5	0.040	0.074
		4.0	0.045	− 0.6	− 0.006	0.085
		28.0	0.312	13.8	0.156	0.107
	80	19.6	0.219	3.8	0.045	0.113
		12.6	0.141	− 0.1	− $26 \times 10^{-6}$	0.087
		9.0	0.101	− 3.2	− 0.034	0.107
		3.0	0.033	− 6.2	− 0.067	0.096
		23.9	0.266	12.9	0.145	0.131
		17.8	0.198	5.8	0.067	0.131
	150	15.3	0.171	4.7	0.054	0.131
		9.3	0.104	− 1.7	− 0.017	0.119
		2.8	0.031	− 7.3	− 0.080	0.101



**Fig. 5.** Hypothesis concerning the usage of  $\text{CO}_2$  and  $\text{HCO}_3^-$  for POC production. Under high external  $[\text{CO}_2]$  and low light intensities (left), POC is built from external  $\text{CO}_2$  which diffuses into the cell. Under low external  $[\text{CO}_2]$  and high light intensities (right), provision of POC production with  $\text{CO}_2$  via diffusion is insufficient. Thus,  $\text{HCO}_3^-$  is taken up as well to provide POC production with  $\text{CO}_2$ . Illustration includes drawing of van der Wal et al. (1983).

#### 4.3. Can $\text{CO}_2$ be accumulated around RubisCO without inorganic carbon species being transported against concentration gradients?

In the cell model,  $\text{CO}_2$  diffuses across plasma membrane and chloroplast envelope into the chloroplast stroma. Inside the alkaline chloroplast stroma, CA-catalysed  $\text{CO}_2$  conversion leads to  $\text{HCO}_3^-$  accumulation. Bicarbonate ions enter the acidic thylakoid lumen through channels and are converted by CA activity to  $\text{CO}_2$ . Inside the thylakoid lumen,  $\text{CO}_2$  is fixed into POC with the rates

measured by Rokitta and Rost (2012) (plus the additional amount that will be lost again to respiration inside the cytosol).

Calculated  $\text{CO}_2$  concentrations inside the thylakoid (Tab. 5) not only exceed external  $\text{CO}_2$  concentrations (Tab. 1) but also the half saturation constant  $K_m$  of RubisCO ( $0.072 \text{ mol} \cdot \text{m}^{-3}$ , Boller et al., 2011). Thus, under LLC, LLHC, and HLHC conditions, the implemented CCM leads to strong  $\text{CO}_2$  accumulation around RubisCO without inorganic carbon species being transported against a concentration gradient.

## 5. Conclusion

A numerical cell model was implemented that proposes an energy-efficient possibility to take up inorganic carbon for POC and PIC production. We find (1) that  $\text{CO}_2$  and  $\text{HCO}_3^-$  do rarely inter-convert during their passage of the cytosol, (2) that  $\text{CO}_2$  for photosynthesis can be taken up via diffusion under high  $\text{CO}_2$  concentrations and nonsaturating light conditions, (3) that a CCM can lead to  $\text{CO}_2$  accumulation around RubisCO without a transport of inorganic carbon species against concentration gradients, and (4) that  $\text{CO}_2$  does not have to leave the cell in order to dispose of  $^{13}\text{CO}_2$  that accumulates intracellularly due to RubisCO's fractionation. Under low external  $\text{CO}_2$  concentrations and high light intensities, the carbon demand of photosynthesis cannot be

**Table 8**

Model inputs.  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  concentrations (unit here:  $\text{mmol} \cdot \text{m}^{-3}$ , unit for implementation:  $\text{mol} \cdot \text{m}^{-3}$ ) in the medium (SW, set constant in the model, measured by Rost et al. (2002)) and the fixation rates of  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  into  $\text{PO}^{12}\text{C}$  and  $\text{PO}^{13}\text{C}$ , respectively ( $R_{\text{PO}^{12}\text{C}}$  and  $R_{\text{PO}^{13}\text{C}}$ , unit here:  $\text{fmol} \cdot \text{cell}^{-1} \cdot \text{h}^{-1}$ , unit for implementation:  $\text{mol} \cdot \text{cell}^{-1} \cdot \text{h}^{-1}$ ). CP - chloroplast; L:D - light:dark cycle (unit: h: h); PFD - photon flux density (unit:  $\mu\text{mol photons} \cdot \text{m}^{-2} \cdot \text{s}^{-1}$ ).

L:D	PFD	SW		CP	
		$^{12}\text{CO}_2$	$^{13}\text{CO}_2$	$R_{\text{PO}^{12}\text{C}}$	$R_{\text{PO}^{13}\text{C}}$
24:0	15	33.4	0.372	19.3	0.212
		30.1	0.334	17.7	0.193
		19.0	0.211	17.7	0.194
		19.7	0.219	18.0	0.198
		13.1	0.145	17.4	0.191
		27.0	0.300	39.8	0.436
	30	23.0	0.256	34.6	0.379
		18.4	0.205	31.3	0.342
		16.7	0.186	33.6	0.367
		11.8	0.131	27.5	0.301
		28.9	0.321	83.6	0.915
		27.0	0.300	68.1	0.745
	80	19.3	0.215	64.6	0.706
		19.1	0.212	56.9	0.622
		11.6	0.129	52.1	0.569
		29.7	0.330	65.9	0.720
		22.4	0.249	58.5	0.640
		17.7	0.197	69.8	0.763
	150	16.6	0.185	62.6	0.685
		11.6	0.129	55.1	0.603
16:8	30	31.7	0.353	28.9	0.320
		23.5	0.262	27.8	0.307
		16.0	0.178	50.3	0.556
		10.8	0.120	37.8	0.418
		5.2	0.058	30.4	0.335
		31.8	0.354	93.8	1.033
	80	23.9	0.266	104.7	1.153
		16.1	0.179	84.6	0.934
		12.4	0.138	81.2	0.894
		5.4	0.061	60.4	0.666
		26.9	0.299	72.8	0.800
		21.1	0.234	79.2	0.871
	150	18.2	0.202	70.2	0.772
		12.3	0.136	72.5	0.797
		5.5	0.062	67.1	0.739

covered by  $\text{CO}_2$  diffusion only. Thus, external  $\text{HCO}_3^-$  has to be used as an additional carbon source which is well in line with the experimental findings of Rost et al. (2002); Bach et al. (2013), and Kottmeier et al..

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### Appendix A

The four differential equations for the stable carbon isotope model depicted in Figure 4 read

$$\begin{cases} d[^{12}\text{CO}_2]^{\text{CS}}/dt \\ d[^{13}\text{CO}_2]^{\text{CS}}/dt \\ d[^{12}\text{CO}_2]^{\text{CP}}/dt \\ d[^{13}\text{CO}_2]^{\text{CP}}/dt \end{cases} = \begin{cases} R_{12\text{CO}_2,\text{PM}}^{\text{CS}} - R_{12\text{CO}_2,\text{CPm}}^{\text{CS}} \\ R_{13\text{CO}_2,\text{PM}}^{\text{CS}} - R_{13\text{CO}_2,\text{CPm}}^{\text{CS}} \\ R_{12\text{CO}_2,\text{CPm}}^{\text{CP}} - R_{\text{PO}^{12}\text{C}} \\ R_{13\text{CO}_2,\text{CPm}}^{\text{CP}} - R_{\text{PO}^{13}\text{C}} \end{cases} \quad (44)$$

where  $[^{12}\text{CO}_2]^{\text{CS}}$  and  $[^{13}\text{CO}_2]^{\text{CS}}$  stand for the concentrations of  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  inside the cytosol and  $[^{12}\text{CO}_2]^{\text{CP}}$  and  $[^{13}\text{CO}_2]^{\text{CP}}$  for the concentrations of  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  inside the bulk chloroplast.

$R_{12\text{CO}_2,\text{PM}}^{\text{CS}}$  and  $R_{13\text{CO}_2,\text{PM}}^{\text{CS}}$  denote the rates by which the concentrations of  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  inside the cytosol change due to aquaporin-based diffusion across the plasma membrane.  $R_{12\text{CO}_2,\text{CPm}}^{\text{CP}}$  and  $R_{13\text{CO}_2,\text{CPm}}^{\text{CP}}$  give the corresponding rates for the bulk chloroplast.

$$R_{12\text{CO}_2,\text{PM}}^{\text{CS}} = \gamma_{\text{CO}_2,\text{PM}} \cdot \frac{A^{\text{CS}}}{V^{\text{CS}}} \cdot ([^{12}\text{CO}_2]^{\text{SW}} - [^{12}\text{CO}_2]^{\text{CS}}) \quad (45)$$

$$R_{13\text{CO}_2,\text{PM}}^{\text{CS}} = \gamma_{\text{CO}_2,\text{PM}} \cdot \frac{A^{\text{CS}}}{V^{\text{CS}}} \cdot ([^{13}\text{CO}_2]^{\text{SW}} - [^{13}\text{CO}_2]^{\text{CS}}) \quad (46)$$

$\gamma_{\text{CO}_2,\text{PM}}$  is the aquaporin-based permeability coefficient for  $\text{CO}_2$  (Tab. 4) and  $[^{12}\text{CO}_2]^{\text{SW}}$  and  $[^{13}\text{CO}_2]^{\text{SW}}$  are given by Rost et al. (2002) (Tab. 8).  $A^{\text{CS}}$  and  $V^{\text{CS}}$  denote plasma membrane amount and cytosol volume, respectively (Tab. 3).

$$R_{12\text{CO}_2,\text{PM}}^{\text{CP}} = R_{12\text{CO}_2,\text{PM}}^{\text{CS}} \cdot \frac{V^{\text{CS}}}{V^{\text{CP}}} \quad (47)$$

$$R_{13\text{CO}_2,\text{PM}}^{\text{CP}} = R_{13\text{CO}_2,\text{PM}}^{\text{CS}} \cdot \frac{V^{\text{CS}}}{V^{\text{CP}}} \quad (48)$$

where  $V^{\text{CP}}$  (Tab. 3) gives the volume of the bulk chloroplast.

$R_{12\text{CO}_2,\text{CPm}}^{\text{CP}}$  and  $R_{13\text{CO}_2,\text{CPm}}^{\text{CP}}$  indicate the rates by which the concentrations of  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  inside the bulk chloroplast change due to aquaporin-based diffusion across the chloroplast envelope.  $R_{12\text{CO}_2,\text{CPm}}^{\text{CS}}$  and  $R_{13\text{CO}_2,\text{CPm}}^{\text{CS}}$  give the corresponding rates for the cytosol.

$$R_{12\text{CO}_2,\text{CPm}}^{\text{CP}} = \gamma_{\text{CO}_2,\text{CPm}} \cdot \frac{A^{\text{CP}}}{V^{\text{CP}}} \cdot ([^{12}\text{CO}_2]^{\text{CS}} - [^{12}\text{CO}_2]^{\text{CP}}) \quad (49)$$

$$R_{13\text{CO}_2,\text{CPm}}^{\text{CP}} = \gamma_{\text{CO}_2,\text{CPm}} \cdot \frac{A^{\text{CP}}}{V^{\text{CP}}} \cdot ([^{13}\text{CO}_2]^{\text{CS}} - [^{13}\text{CO}_2]^{\text{CP}}) \quad (50)$$

$$R_{12\text{CO}_2,\text{CPm}}^{\text{CS}} = R_{12\text{CO}_2,\text{CPm}}^{\text{CP}} \cdot \frac{V^{\text{CP}}}{V^{\text{CS}}} \quad (51)$$

$$R_{13\text{CO}_2,\text{CPm}}^{\text{CS}} = R_{13\text{CO}_2,\text{CPm}}^{\text{CP}} \cdot \frac{V^{\text{CP}}}{V^{\text{CS}}} \quad (52)$$

$\gamma_{\text{CO}_2,\text{CPm}}$  gives the  $\text{CO}_2$  permeability coefficient of the chloroplast envelope (Tab. 4).  $A^{\text{CP}}$  denotes the surface area of the chloroplast.

$R_{\text{PO}^{12}\text{C}}$  and  $R_{\text{PO}^{13}\text{C}}$  give the fixation rates of  $^{12}\text{CO}_2$  and  $^{13}\text{CO}_2$  into  $\text{PO}^{12}\text{C}$  and  $\text{PO}^{13}\text{C}$ , respectively, that were calculated from the data of Rost et al. (2002).

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